Multiscale techniques for the extraction of quantitative information from chlorophyll and SST maps

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Overview

. Velocity PDF.

. Singularity Spectrum.

- . Examples of application.
- . Scale analysis.

. Scale invariance properties.

. Conclusions.

Velocity Probability Density Functions



Non-Gaussian character of PDF: ocean features, e.g., eddies.
Vortices: key role in ocean dynamics.

(J. Isern-Fontanet et al., 2006, JPO)



Velocity Probability Density Functions



FIG. 1. Observed velocity pdfs for three 2.5° boxes. Open circles are from the South Atlantic, black triangles from the South Pacific Ocean, and gray diamonds from the energetically varying Malvinas Current in the South Atlantic, an exponential distribution. Solid lines show best fit Gaussian or exponential pdfs.

. Global scale: spatial inhomogeneity of EKE.

. PDF appearance changes: domain size, lat. and lon.

. S. G. Llewellyn Smith et al., 1998, Phys. Rev. Letters.



distribution worrsubjected to scalar, scale, r/24edign daily.



* DOMINANT ADVECTION: scalars - tracers of ocean dynamics.



Singularity exponents map from MW SST

Sequence of SE maps: TIW



. Flow agitation state transmitted to SE: underlying physics of FDT.

. Reliable approximation of REAL FLOWS behavior. . NEW PERSPECTIVE of turbulent flow events.



. A. Turiel, J. Sole, V. Nieves, J. Ballabrera-Poy, and E. Garcia-Ladona, 2007, RSE. . A. Turiel, J. Sole, V. Nieves, and E. Garcia-Ladona, 2007, IGARSS.

SERESIANAPAR

GULF

STREAM

Comparison between SE from different variables

SE from different advected tracers must coincide:



. Application standpoint: SST can be used as a PROXY of CC.

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. The Singularity Spectrum is the same FOR ANY SCALAR.

. Singularity Spectrum:

$$D(h) = d - \frac{\log\left(\frac{\rho(h)}{\rho(h_1)}\right)}{\log r}$$

. ρ(h) - empirical histogram.
. r - resolution scale.
(Turiel et al., 2006, JCP.)



. V. Nieves et al., 2007, Geophys. Res. Lett.





- . CLS, AVISO server.
- . T/P improved by ERS.
- . Subtract 4-year mean: systematic deviations.
- . Lanczos filter +
- long wavelength errors.
- * Res.1/3 deg., 10 days.

- . GlobColour, ESA.
- . Level-3 ocean color:
- MERIS+MODIS+SeaWiFS.
- . GSM: merging method.
- * Res. 1/24 deg., daily.

- . Rem. Sen. Syst., NASA.
- . Interpol. techniques: missing data resolved.
- . Nearly global coverage.
- * Res. 1/4 deg., daily.

Scale analysis of the Singularity Spectrum

. Scale analysis of SS allows detecting interpolation artifacts.



Detection of interpolation artifacts

OSTIA: Operational Sea Surface Temperature and Sea Ice Analysis



AMSRE-E:



- . Satellite (GHRSST) + in-situ data.
- . Optimal Interpolation Analysis:
- AMRSE + TMI + Pathfinder + MODIS.
- * Res. 1/20 deg., daily.

INTERPOLATED DATA: NOMINAL RES. X 5 = DATA

- . Rem. Sen. Syst., NASA.
- . Capability through clouds.
- * Res. 1/4 deg., daily.

Detection of interpolation artifacts

Aqua - MODIS ocean color sensor: daytime SST, res. 1/24 deg., daily.



Conclusions

. Local degree of correspondence between SE of different satellite data.

. SE of adequate scalar vs. other maps with gaps - inference.

- Future merging / interpolation.
- Improvement of ocean modelling, bio-physical, and climatic studies.

. SS allows describing scaling properties: scale-invariance.

- . Interpolated satellite products show some limitations.
- . This can be overcome increasing data statistics and resolution.

http://www.icm.csic.es/oce/people/vnieves/

THANK YOU FOR YOUR ATTENTION!







Inferring missing data

. Cascade relation between is selected for the gap area. Every father coefficient is inferred from its 4 children.



$$\alpha_{c} = \eta \cdot \alpha_{p} \Longrightarrow \log_{2} |\alpha_{c}| = \log_{2} |\eta| + \log_{2} |\alpha_{p}|$$

$$\log_2 |\alpha_{C_1}| - \log_2 |\alpha_P| = \log_2 |\eta_1|$$

$$\log_2 |\alpha_{C_2}| - \log_2 |\alpha_P| = \log_2 |\eta_2|$$

$$\log_2 |\alpha_{C_3}| - \log_2 |\alpha_P| = \log_2 |\eta_3|$$

$$\log_2 |\alpha_{C_4}| - \log_2 |\alpha_P| = \log_2 |\eta_4|$$



E transfers: cascade process

OSTIA:

is this hypothesis particularly well be haved under $\sum_{S(\mathcal{X})}^{J} \sum_{j=1}^{J} \sum_{\vec{k}} \alpha_{rj\vec{k}} \psi_{rj\vec{k}}(\vec{x})$ atmospheric forcing / specific ocean conditions $\sum_{j=1,2,3}^{J} \sum_{j=1}^{J} \sum_{\vec{k}} \alpha_{rj\vec{k}} \psi_{rj\vec{k}}(\vec{x})$

. Is the E-transfer significantly relevant at a $\alpha_{rj\vec{k}} \doteq \eta_{rj\vec{k}} \alpha_{r,j+1,\left[\frac{\vec{k}}{2}\right]}$

Ratio between scales: there is a background E dissipation.



$$\eta_{rj\vec{k}} = \frac{\alpha_{rj\vec{k}}}{\alpha_{r,j+1,\left[\frac{\vec{k}}{2}\right]}}$$

. O. Pont et al., in preparation.